# Variation of Nongravitational Parameters for Comet Encke as a Result of its Decay 

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#### Abstract

An assumption is made that variations of nongravitational parameters of comet Encke in Marsden's model are caused by its secular decay. To account for decrease of nongravitational parameters three variants of physical processes are proposed: 1) deposition of a substantial nonvolatile mass not impeding the sublimation; 2) generation of a little-massive mantle shrinking the effective area linearly; 3) generation of a little-massive mantle shrinking the area exponentially. The corresponding equations are derived. The adequacy of the first two models over more than 150 years is shown.


Many models [1] explain the behavior of Marsden's parameters [2] of comet Encke [3] by the rotation pole precession of the spotty nucleus as proposed by Whipple and Sekanina [4]. Since the comet decay is not considered, they are not self-consistent. Other shortcomings are reviewed by Chernetenko [5].

In the present work we examine three models for secular decay of comet Encke assuming a constant shape of the nucleus. (Keeping nearly constant shape of a cometary nucleus was confirmed by Medvedev [6].) Since not all the mass is ejected in the same direction, Meshcersky's equation for the reactive force acting on the comet nucleus of mass $m$ is given by

$$
\begin{equation*}
\vec{F}_{r}=\lambda \vec{u} \frac{d m}{d t}, \tag{1}
\end{equation*}
$$

where $\lambda$ is the anisotropy factor and $\vec{u}$ is the velocity of the matter escaping from the nucleus in the orbital coordinate system; in Marsden's model [2]

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both being considered as constants. The number of particles in mass $m$ is given by

$$
\begin{equation*}
N=m N_{A} / M \tag{2}
\end{equation*}
$$

where $N_{A}$ is the Avogadro number and $M$ is the mean molecular mass. In Marsden's model the number of ejected particles from a unit pure area during unit time is given by

$$
\begin{equation*}
-\frac{1}{\beta S} \frac{d N}{d t} \equiv Z=Z_{0} g(r(t)), \tag{3}
\end{equation*}
$$

where $S$ is the geometrical area, $0 \leq \beta \leq 1$ is the ratio of the effective area to geometrical area, $r(t)$ is the heliocentric distance (in AU ), and

$$
\begin{equation*}
g(r)=0.111262 \cdot 10^{-8}(r / 2.808)^{-2.15}\left(1+(r / 2.808)^{5.093}\right)^{-4.6142} \tag{4}
\end{equation*}
$$

In preliminary calculations one can use $\langle g(r(t))\rangle=g(a, \quad e)$ depending on the size and shape of the orbit. One obtains the acceleration in orbital coordinates ( $i=1 ; 2 ; 3$ are radial, transverse, and normal directions):

$$
\begin{equation*}
w_{i}=-\frac{\lambda u_{i} M Z_{0} \beta S}{N_{A} m} g(r) . \tag{5}
\end{equation*}
$$

By definition, one finds Marsden's parameters (units are AU/(104 days)2):

$$
\begin{equation*}
A_{i}=-\frac{\lambda u_{i} M Z_{0} \beta S}{N_{A} m} \tag{6}
\end{equation*}
$$

1. Assume a substantial nonvolatile mass is deposited. To continue this deposition, it should not impede the sublimation ( $\beta=1$ ). Before Style II Marsden's model was introduced, a similar case was considered by Sekanina [7] but it was not developed. If the nucleus shape is constant, then

$$
\begin{equation*}
\frac{S}{m_{\mathrm{ice}}}=\frac{\varphi}{\rho_{\mathrm{ice}} R}, \tag{7}
\end{equation*}
$$

where $\varphi$ depends on the shape $\left(\varphi=3\right.$ for a sphere), $\rho_{\text {ice }}$ and $m_{\text {ice }}$ are the density and the mass of the ice, $R$ is its mean radius defined as

$$
\begin{equation*}
R=\sqrt[3]{\frac{3 m_{\mathrm{ice}}}{4 \pi \rho_{\mathrm{ice}}}} \tag{8}
\end{equation*}
$$

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Analogically $R_{\text {end }}$ is defined using the nonvolatile mass $m_{\text {end }}=m-m_{\text {ice }}$ :

$$
\begin{equation*}
R_{\mathrm{end}}=\sqrt[3]{\frac{3 m_{\mathrm{end}}}{4 \pi \rho_{\mathrm{end}}}} . \tag{9}
\end{equation*}
$$

Hence Marsden's parameters and their variations due to the ablation are

$$
\begin{gather*}
A_{i}=-\frac{\lambda u_{i} M Z_{0} \varphi}{N_{A} \rho_{\mathrm{ice}}} \frac{R^{2}}{R^{3}+\left(\rho_{\mathrm{end}} R_{\mathrm{end}}^{3} / \rho_{\mathrm{ice}}\right)},  \tag{10}\\
\frac{d A_{i}}{d R}=-\frac{\lambda u_{i} M Z_{0} \varphi}{N_{A} \rho_{\mathrm{ice}}} \frac{2 R\left(R^{3}+\left(\rho_{\mathrm{end}} R_{\mathrm{end}}^{3 /} \rho_{\mathrm{ice}}\right)\right)-3 R^{4}}{\left(R^{3}+\left(\rho_{\mathrm{end}} R_{\mathrm{end}}^{3} / \rho_{\mathrm{ice}}\right)\right)^{2}} . \tag{11}
\end{gather*}
$$

And also their combination is

$$
\begin{equation*}
\frac{d A_{i}}{A_{i} d R}=\frac{2}{R}-\frac{3 R^{2}}{R^{3}+\left(\rho_{\mathrm{end}} R_{\mathrm{end}}^{3} / \rho_{\mathrm{ice}}\right)} \tag{12}
\end{equation*}
$$

Combining (2), (3), (7), and (8) if $\beta=1$, one obtains

$$
\begin{equation*}
\frac{d R}{d t}=-\frac{\varphi M Z_{0}}{3 N_{A} \rho_{\mathrm{ice}}} g(r) \tag{13}
\end{equation*}
$$

Designating $\sqrt[3]{\frac{\rho_{\mathrm{ice}}}{\rho_{\text {end }} R_{\mathrm{end}}{ }^{3}}} R=\chi \geq 0$ and $\frac{\phi M Z_{0}}{3 N_{A} \rho_{\mathrm{ice}}{ }^{2 / 3} \rho_{\text {end }}{ }^{1 / 3} R_{\text {end }}}=\alpha \geq 0$ in (12) and (13), one writes the final set

$$
\left\{\begin{align*}
\frac{d A_{i}}{d \chi} & =A_{i}\left(\frac{2}{\chi}-\frac{3 \chi^{2}}{1+\chi^{3}}\right)  \tag{14}\\
\frac{d \chi}{d t} & =-\alpha g(r(t)) \\
\alpha & =\text { const. }
\end{align*}\right.
$$

Table 1. Parameters of solution in Fig. 1 for beginning and ending dates of model

| Year | $A_{2}$ | $\alpha$, day $^{-1}$ | $\chi$ |
| :---: | :---: | :---: | :---: |
| 1786 | -0.0461 | $3.24 \cdot 10^{-5}$ | 0.945 |
| 2032 | 0 |  | 0 |

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Fig. 1. Formal solution of set (14) on linear and logarithmic scales.
The formal (because input data are $A_{2}$ with mean errors from [2, 3, 8] and other sources, not astrometry) solution of (14) for $A_{2}$ (which is much more accurate than both $A_{1}$ and $A_{3}$ ) is in Fig. 1, its parameters are in Tab. 1. Here and below elements $a, e$, are assumed as equal to their mean values. The accuracy of $A_{2}$ increased with time when its value decreased, thus the graph on the logarithmic scale is a better representation for the weighted accuracy. One can see that the model reproduces $A_{2}$ adequately over more than 170 years and predicts the total decay about 2022.
2. Suppose a little-massive mantle is generated shrinking the effective area linearly with the thickness. This is equivalent to Shul'man's assumption that the mantle formation is completed when the volume containing the covering area equal to the nucleus area has sublimated [9]. In analogy to (7) one has

$$
\begin{equation*}
\frac{S}{m}=\frac{\varphi}{\rho R}, \tag{15}
\end{equation*}
$$

where $\rho$ is the nucleus density (assumed uniform), $R$ is the mean radius:

$$
\begin{equation*}
R=\sqrt[3]{\frac{3 m}{4 \pi \rho}} \tag{16}
\end{equation*}
$$

Hence Marsden's parameters and their variations due to the ablation are

$$
\begin{align*}
& A_{i}=-\frac{\lambda u_{i} M Z_{0} \varphi \beta}{N_{A} \rho R}  \tag{17}\\
& \frac{d A_{i}}{d R}=-\frac{\lambda u_{i} M Z_{0} \varphi}{N_{A} \rho} \frac{d(\beta / R)}{d R} . \tag{18}
\end{align*}
$$

And also their combination is

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$$
\begin{equation*}
\frac{d A_{i}}{A_{i} d R}=\frac{R}{\beta} \frac{d(\beta / R)}{d R}=\frac{1}{\beta} \frac{d \beta}{d R}-\frac{1}{R} \tag{19}
\end{equation*}
$$

By the model assumption, one gets

$$
\begin{equation*}
\beta=1-h / h_{0} \tag{20}
\end{equation*}
$$

where $h_{0}$ is the mantle thickness terminating the sublimation,

$$
\begin{equation*}
h=\frac{f\left(m_{\text {begin }}-m\right) / \rho}{S} \tag{21}
\end{equation*}
$$

is its current thickness, $m_{\text {begin }}$ is the initial mass, and $f$ is the nonvolatiles bulk part. Substituting (15) and (16) one has

$$
\begin{equation*}
h=\frac{f}{\varphi} \frac{R_{\mathrm{begin}}{ }^{3}-R^{3}}{R^{2}} . \tag{22}
\end{equation*}
$$

Derivating (20) with respect to $R$, considering (22), and substituting $R_{\text {begin }}$ from them, one obtains

$$
\begin{equation*}
\frac{d \beta}{d R}=\frac{3 f}{\varphi h_{0}}+\frac{2(1-\beta)}{R} . \tag{23}
\end{equation*}
$$

Hence one sets (19) to:

$$
\begin{equation*}
\frac{d A_{i}}{d R}=A_{i}\left(\frac{1}{\beta} \frac{3 f}{\varphi h_{0}}+\frac{2}{\beta R}-\frac{3}{R}\right) . \tag{24}
\end{equation*}
$$

Combining (2), (3), (15), and (16), one obtains

$$
\begin{equation*}
\frac{d R}{d t}=-\frac{\varphi M Z_{0}}{3 N_{A} \rho} \beta g(r) . \tag{25}
\end{equation*}
$$

Designating $\frac{3 f}{\varphi h_{0}} R=\chi$ and $\frac{M Z_{0} f}{N_{A} \rho h_{0}}=\alpha$ in (24), (23), and (25), one writes

$$
\left\{\begin{array}{c}
\frac{d A_{i}}{d \chi}=A_{i}\left(\frac{1}{\beta}+\frac{2}{\beta \chi}-\frac{3}{\chi}\right),  \tag{26}\\
\frac{d \beta}{d \chi}=1+\frac{2(1-\beta)}{\chi} \\
\frac{d \chi}{d t}=-\alpha \beta g(r(t)) \\
\alpha=\text { const. }
\end{array}\right.
$$

The formal solution of (26) for $A_{2}$ is in Fig. 2, its parameters are in Tab. 2. It reproduces $A_{2}$ adequately over more than 150 years and suggests that on discovery the comet surface was nearly pure ice ( $\beta>1$ makes no sense).


Fig. 2. Formal solution of set (26) on linear and logarithmic scales.
Table 2. Parameters of solution in Fig. 2 for beginning and ending dates of model

| Year | $A_{2}$ | $\alpha$, day $^{-1}$ | $\beta$ | $\chi$ |
| :--- | :--- | :---: | :---: | :---: |
| 1786 | -0.0372 | $3.32 \cdot 10^{-5}$ | 0.99 | 0.637 |
| 2032 | -0.000455 |  | 0.00547 | 0.288 |

3. A case of the generation of a little-massive mantle shrinking the effective area exponentially was also considered. It was shown that this model represents the variation of $A_{2}$ only qualitatively and is of no interest.

Solutions obtained in the present work leave significant offsets. Causes of offsets may be not only in model assumptions, but also out of them. These are the solution procedure formality, the cometary stochasticity, and accidental errors in $A_{2}$ multiplied by the correlation with little-significant $A_{1}$.

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